Rossby wave breaking and extreme precipitation events in the central and eastern United States

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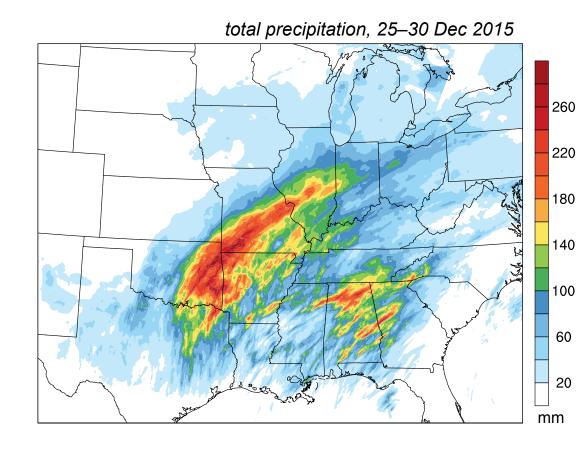
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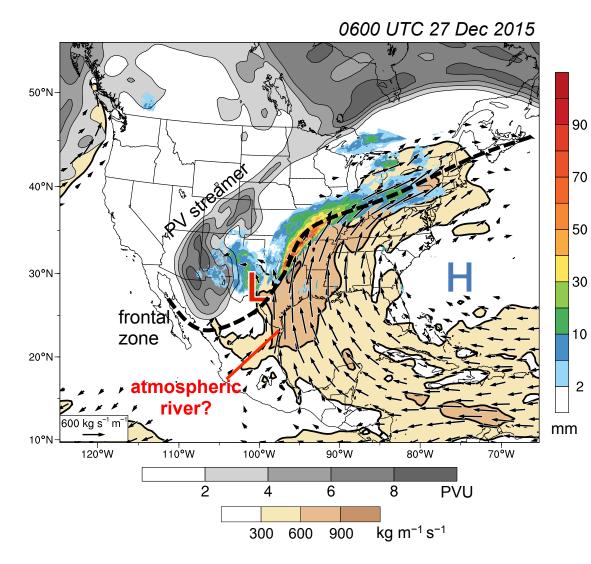
Motivation

- Widespread extreme precipitation events (EPEs) in central and eastern U.S. during cool season can result in highimpact flooding
- Evidence from case studies indicates widespread EPEs occur in conjunction with baroclinic Rossby wave breaking (RWB)
- Climatological and dynamical linkages between RWB and EPEs in U.S. have not yet been examined



Motivation

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- Climatological and dynamical linkages between RWB and EPEs in U.S. have not yet been examined



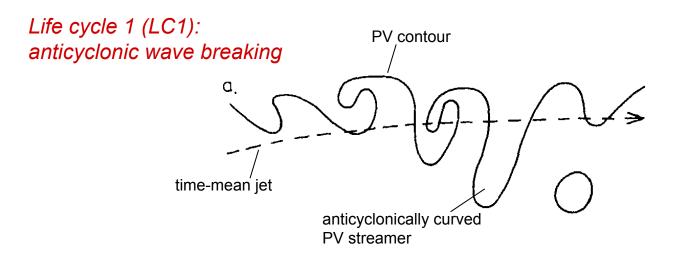
320-K PV (PVU, gray shading), 1000–200-hPa IVT (kg m⁻¹ s⁻¹, vectors and tan shading) 6-h Stage-IV precip (mm, color shading)

Science questions

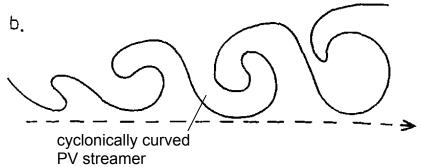
- 1. Do widespread EPEs in the central/eastern U.S. preferentially occur in conjunction with RWB?
- 2. How does RWB supply the ingredients for EPEs?

This presentation will focus on the nexus between RWB and moisture transport linked to EPEs.

RWB: An aspect of baroclinic wave life cycles



Life cycle 2 (LC2): cyclonic wave breaking



Adapted from Fig. 12 in Thorncroft et al. (1993)

baroclinic wave life cycles & Rossby wave breaking

strong moisture transport (i.e., ARs) & dynamical forcing for ascent

heavy precipitation

ARs form as an aspect of the dynamical evolution of baroclinic waves that establishes favorable conditions for heavy precipitation

Data and methods

- Use 24-h (ending 1200 UTC) 0.25° gauge-based precipitation analyses for 1979–2015 from NOAA CPC Unified Precipitation Dataset
- Define top 5% (299) of days with ≥1 extreme precip value in domain as widespread EPEs
- Retain only days during Sep–May without a tropical cyclone in domain
- Consider consecutive days as one event; retain only largest-scale day for statistical analysis
- Examine final sample of 201 widespread EPEs
- Define t₀ as start time of 24-h period of EPE

Climatology of widespread EPEs

99th percentile of 24-h precipitation for all days with >0 mm

50°N

precipitation

domain

35

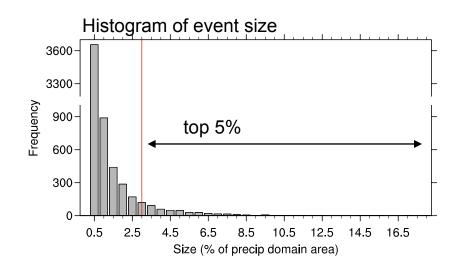
45

55

65

15

25



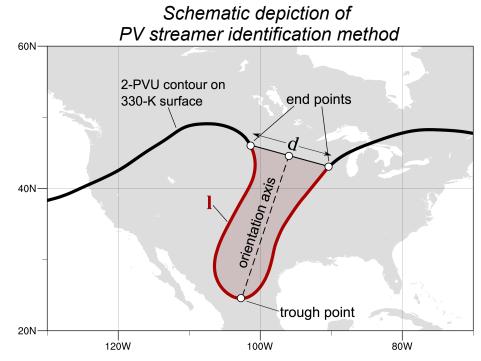
75 mm

Data and methods

Adaptation of method from Wernli and Sprenger (2007)

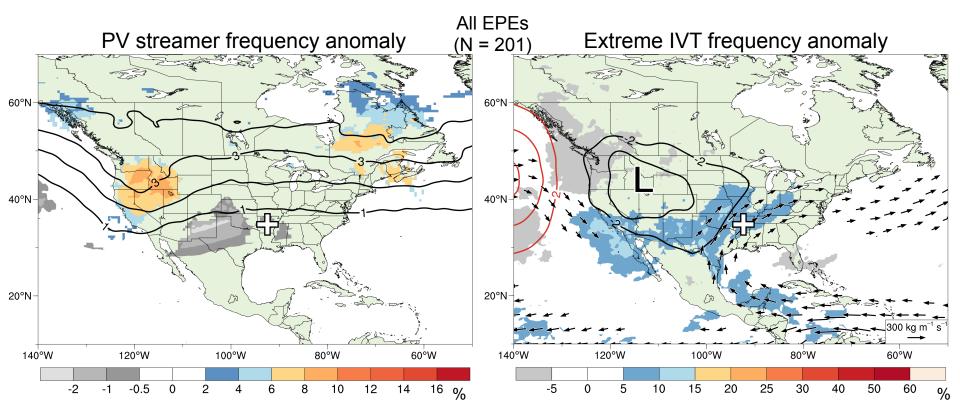
- Potential vorticity (PV) streamers are manifestations of RWB
- PV streamers identified on 2-PVU contour on 310-, 320-, and 330-K surfaces in 0.5° NCEP CFSR
- PV streamers identified as pairs of points along 2-PVU contour separated by distance d < 1000 km and by contour length l > 3000 km
- Classify streamers based on orientation angle relative to meridional baseline through midpoint of interval d:
 - > 15°: anticyclonic (LC1) wave breaking
 - < −15°: cyclonic (LC2) wave breaking
 - o all others: "meridional"

Climatology of breaking Rossby waves



Adapted from Fig. 1 in Wernli and Sprenger (2007)

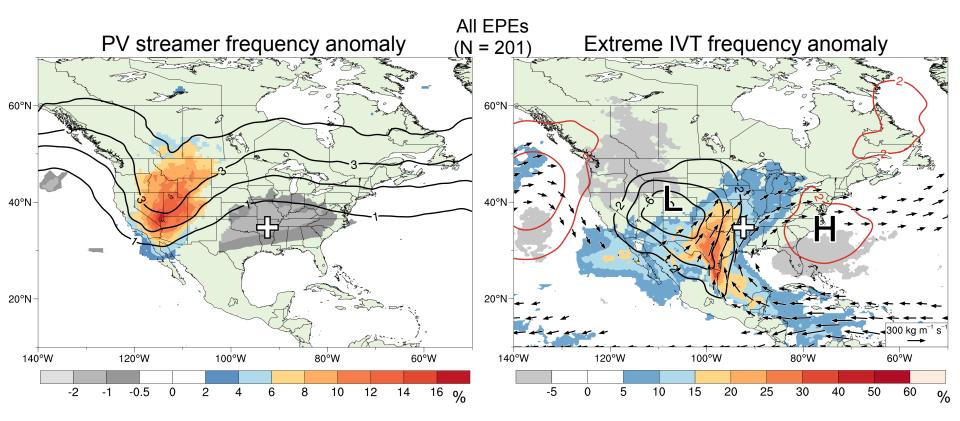
 $t_0 - 48 h$



320-K PV streamer frequency displayed as an anomaly relative to the climatological frequency (%, shading; only statistically significant values shown) and composite 320-K PV (PVU, black)

Frequency of >90th percentile IVT displayed as an anomaly relative to the climatological frequency (%, shading; only statistically significant values shown), and composite IVT vectors (kg m⁻¹ s⁻¹) and SLP anomaly (hPa, negative in black; positive in red)

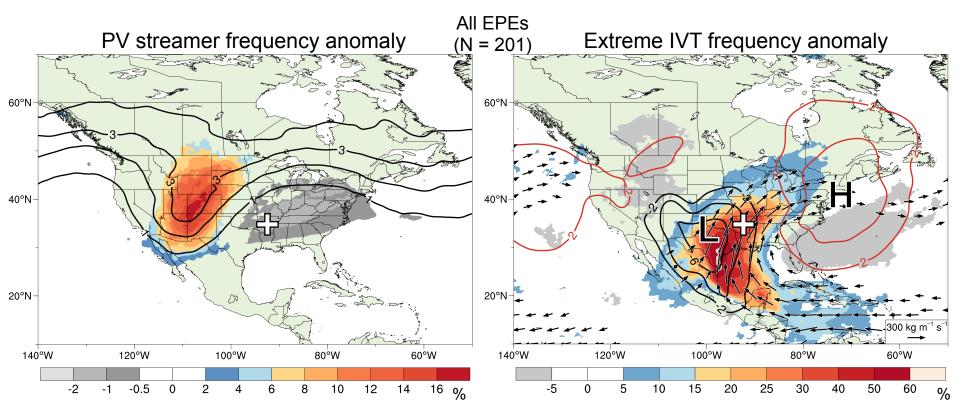
 $t_0 - 24 h$



320-K PV streamer frequency displayed as an anomaly relative to the climatological frequency (%, shading; only statistically significant values shown) and composite 320-K PV (PVU, black)

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 $t_0 - 0 h$

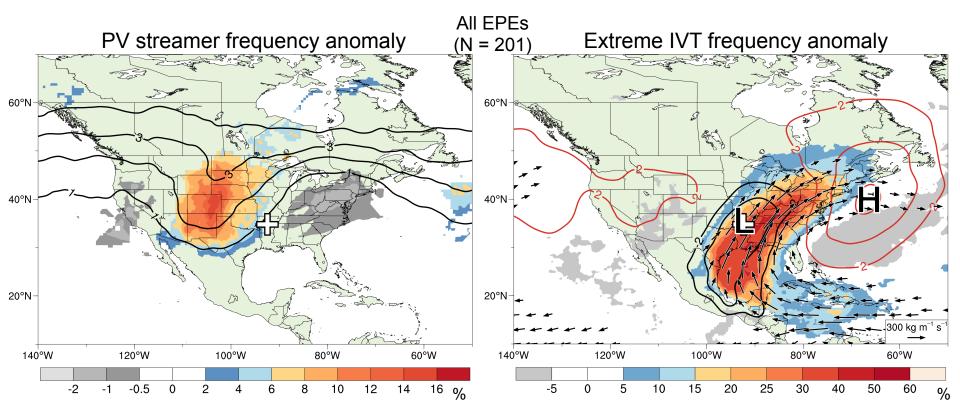


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RWB and extreme water vapor transport linked to EPEs

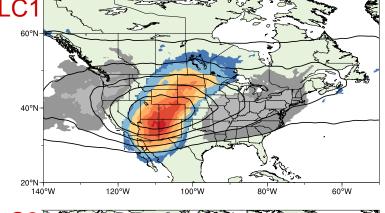
 $t_0 + 24 h$

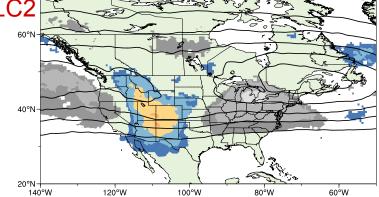


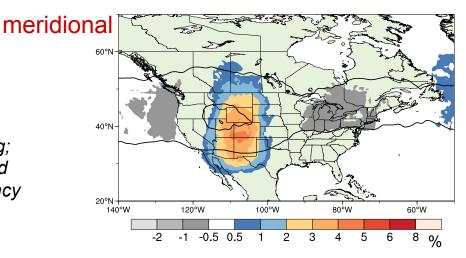
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Preferred regions for PV streamer occurrence associated with EPEs



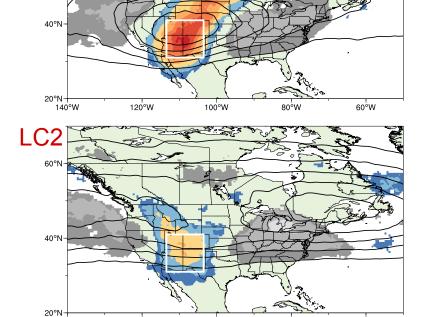


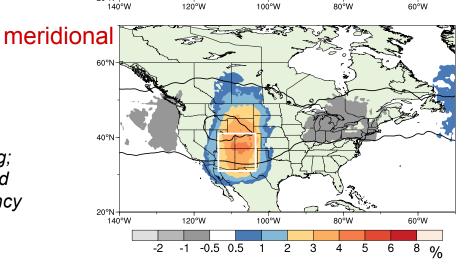


320-K PV streamer frequency anomaly (%, shading; only statistically significant values shown) averaged between t_0 – 12 h and t_0 + 12; climatological frequency contoured in black every 0.5%

Approach

- Identify streamers that overlap ¼ of area of 10° × 10° box
- Consider EPE linked to RWB if streamer identified within 24-h period centered on t₀





320-K PV streamer frequency anomaly (%, shading; only statistically significant values shown) averaged between t_0 – 12 h and t_0 + 12; climatological frequency contoured in black every 0.5%

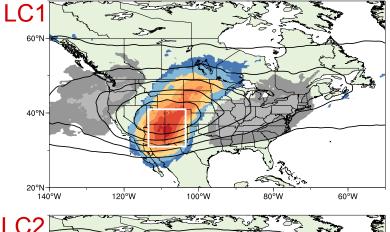
Approach

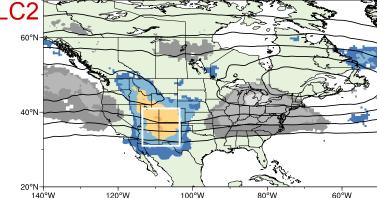
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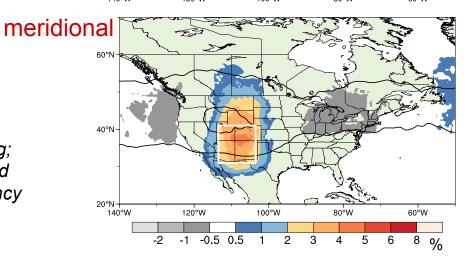
Results for 320-K PV streamers

- ~48% (97 of 201) of EPEs linked to RWB
 - LC1: ~47% (46)
 - o LC2: ~26% (25)
 - meridional: ~27% (26)

320-K PV streamer frequency anomaly (%, shading; only statistically significant values shown) averaged between t_0 – 12 h and t_0 + 12; climatological frequency contoured in black every 0.5%





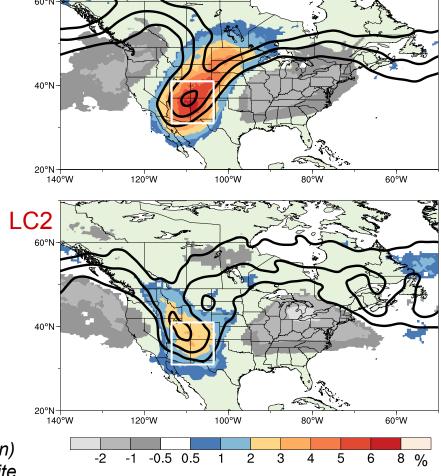


Results when analysis repeated to include PV streamers identified on at least one of three isentropic surfaces (i.e., 310, 320, 330 K)

- ~76% (153 of 201) of EPEs linked to RWB
 - **LC1**: ~49% (75)
 - o **LC2:** ~23.5% (36)
 - meridional: ~27.5% (42)

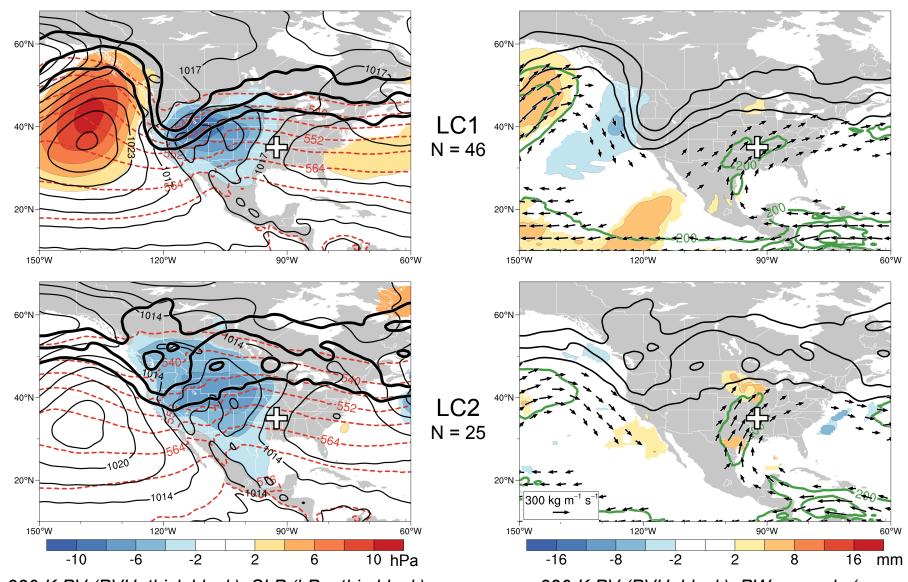
Approach:

- Construct composites for EPEs linked to PV streamers identified on 320-K surface
- Examine only LC1 and LC2 cases to highlight distinct EPE scenarios



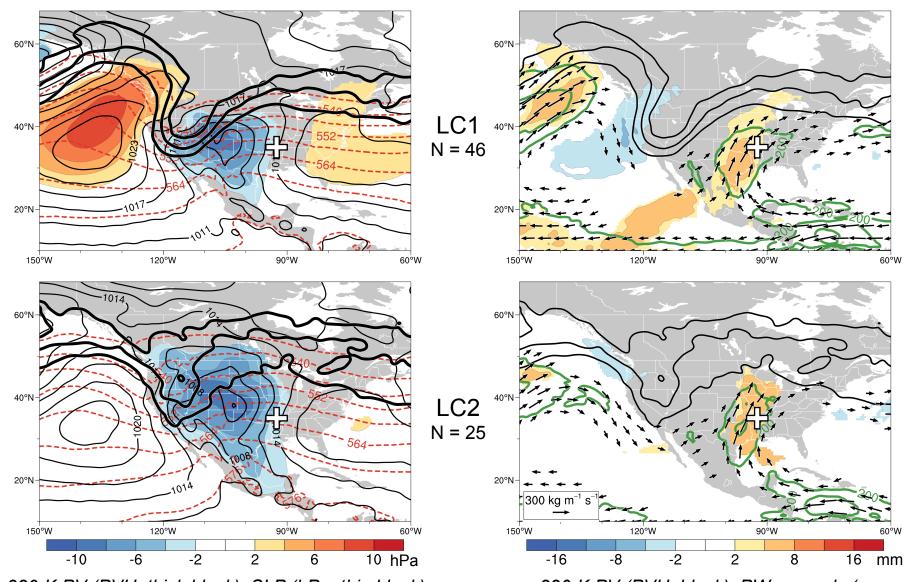
LC₁

320-K PV streamer frequency anomaly (%, shading; only statistically significant values shown) averaged between t_0 – 12 h and t_0 + 12; composite 320-K PV (PVU, black) at t_0 for PV streamers overlapping box



320-K PV (PVU, thick black), SLP (hPa, thin black), SLP anomaly (hPa, shading; only statistically significant values shown), 1000–500-hPa thickness (dam, red)

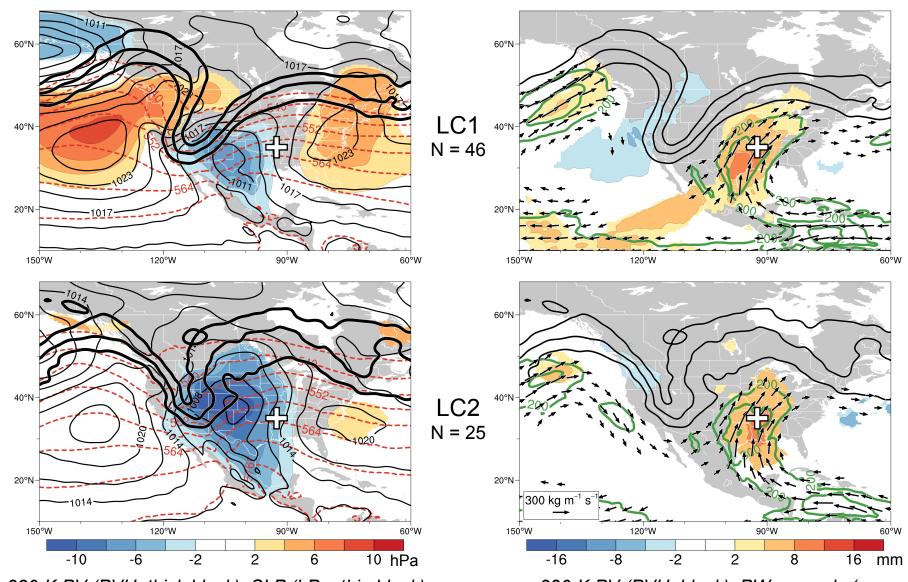
320-K PV (PVU, black), PW anomaly (mm, shading; only statistically significant values shown), IVT (kg m⁻¹ s⁻¹, green and vectors)



320-K PV (PVU, thick black), SLP (hPa, thin black), SLP anomaly (hPa, shading; only statistically significant values shown), 1000–500-hPa thickness (dam, red)

320-K PV (PVU, black), PW anomaly (mm, shading; only statistically significant values shown), IVT (kg m⁻¹ s⁻¹, green and vectors)

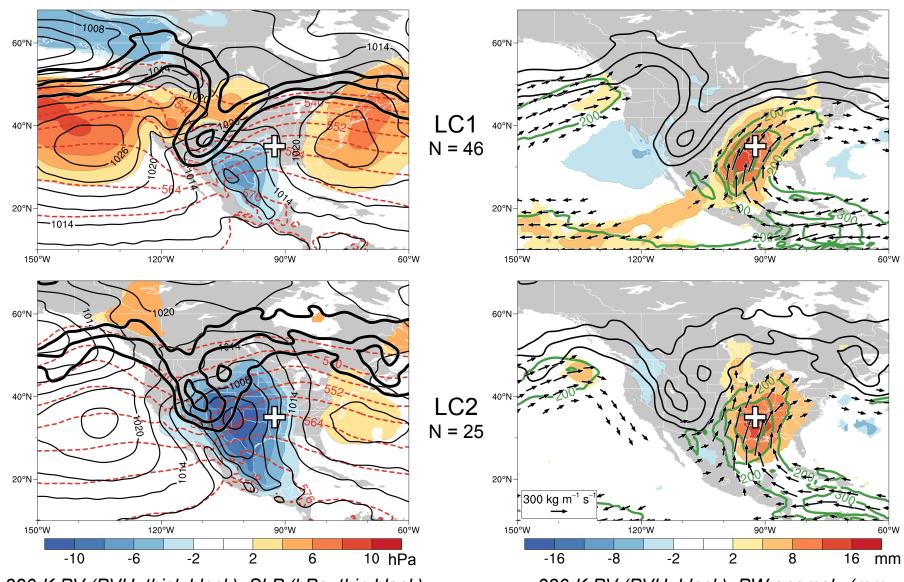
 $t_0 - 12 h$



320-K PV (PVU, thick black), SLP (hPa, thin black), SLP anomaly (hPa, shading; only statistically significant values shown), 1000–500-hPa thickness (dam, red)

320-K PV (PVU, black), PW anomaly (mm, shading; only statistically significant values shown), IVT (kg m⁻¹ s⁻¹, green and vectors)

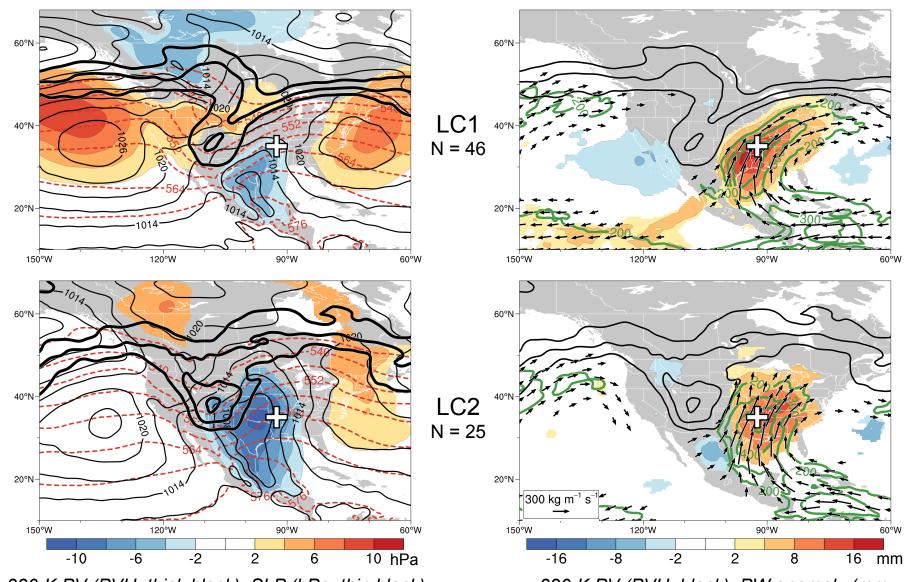
 $t_0 - 0 h$



320-K PV (PVU, thick black), SLP (hPa, thin black), SLP anomaly (hPa, shading; only statistically significant values shown), 1000–500-hPa thickness (dam, red)

320-K PV (PVU, black), PW anomaly (mm, shading; only statistically significant values shown), IVT (kg m⁻¹ s⁻¹, green and vectors)

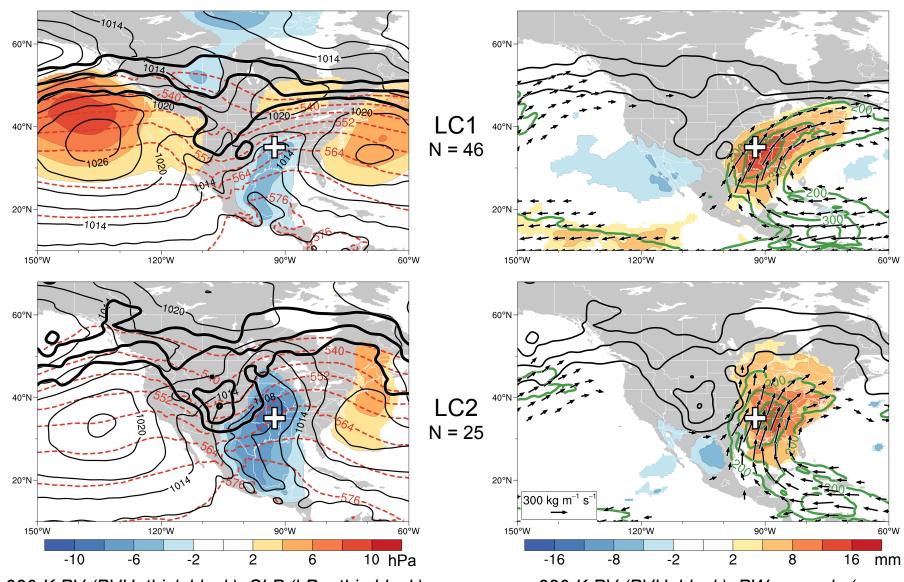
 $t_0 + 12 h$



320-K PV (PVU, thick black), SLP (hPa, thin black), SLP anomaly (hPa, shading; only statistically significant values shown), 1000–500-hPa thickness (dam, red)

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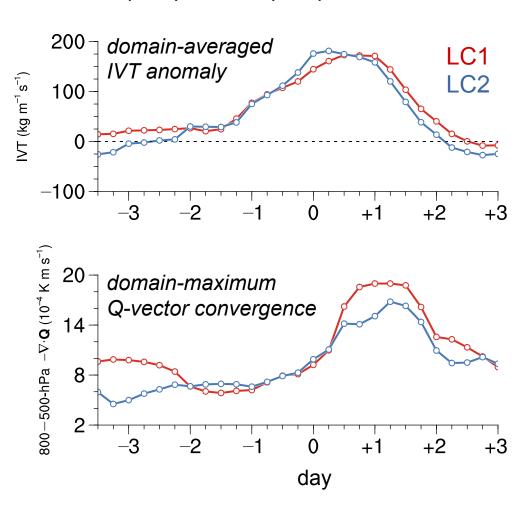
 $t_0 + 24 h$



320-K PV (PVU, thick black), SLP (hPa, thin black), SLP anomaly (hPa, shading; only statistically significant values shown), 1000–500-hPa thickness (dam, red)

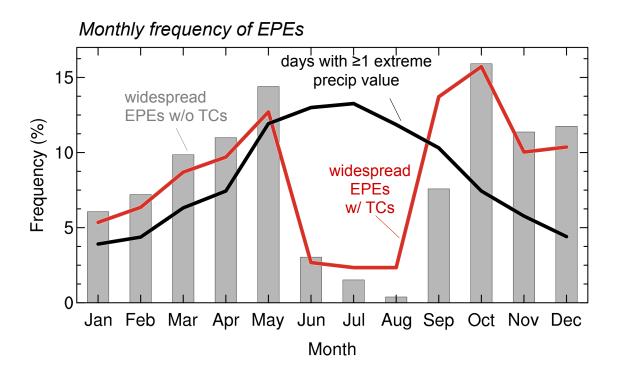
320-K PV (PVU, black), PW anomaly (mm, shading; only statistically significant values shown), IVT (kg m⁻¹ s⁻¹, green and vectors)

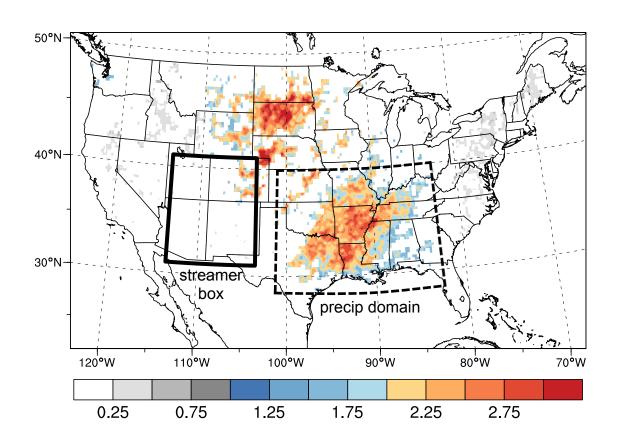
Time series of ingredients for precipitation in precipitation domain



Summary

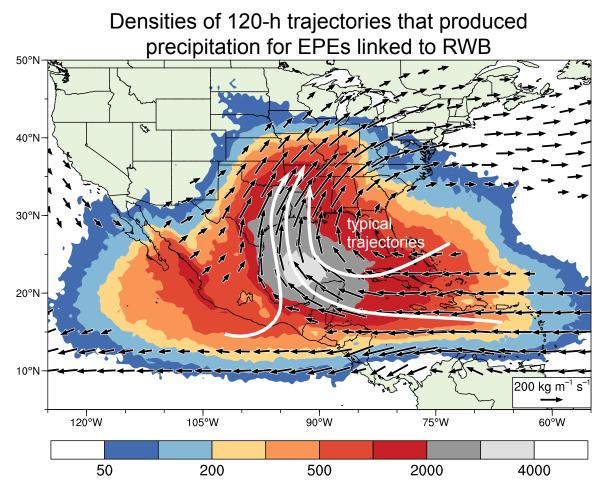
- Climatological and dynamical linkages between RWB and EPEs over portions of the central/eastern U.S. during 1979–2015 examined
- Large majority (~76%) of EPEs examined found to occur in connection with RWB; LC1 dominant relative to LC2
- PV streamers associated with EPEs occur over discrete regions centered over the western U.S.
- RWB linked to formation of high-amplitude, slow-moving wave pattern that establishes persistent corridor of strong water vapor transport (i.e., AR)
- Water vapor transport supports EPE occurrence in presence of dynamical forcing for ascent





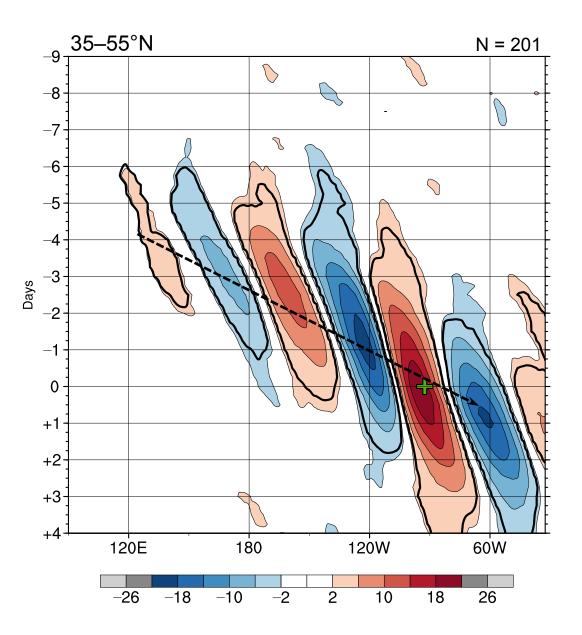
factor of increase in probability of extreme precipitation relative to climatology (shading; only statistically significant values shown) for days during Sep–May on which a streamer overlaps ½ of the area of the box

Lagrangian perspective



Trajectory density for 120-h backward trajectories released during EPE from $5^{\circ} \times 5^{\circ}$ box centered on maximum precipitation location that exhibited >5 g kg⁻¹ decrease in specific humidity in final 24 h; time-mean composite IVT vectors for t_0 – 72 h and t_0 + 24 overlaid

Linkage of EPEs to Rossby waves



Hovmöller of 250-hPa merid. wind anomalies (m s⁻¹, shading), statistical significance at 95% confidence level (black contours)